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SHEAR WAVE MEASUREMENTS IN SHOCK-INDUCED, HIGH-PRESSURE PHASES.*

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Structural phase transformations under shock loading are of considerable interest for understanding the response of solids under nonhydrostatic stresses and at high strain-rates. Examining shock-induced transformations from continuum level measurements is fundamentally constrained by the inability to directly identify microscopic processes, and also by the limited number of material properties that can be directly mensured. The latter limitation can be reduced by measuring both shear and compression waves using Lagrangian gauges in combined, compression and shear loading. The shear wave serves as an important, real-time probe of the shocked state and unloading response. Using results from a recent study of CaCO₃, the unique information obtained from the shear wave speed and the detailed structure of the shear wave are shown to be useful for distinguishing the effects of phase transformations from yielding, as well as in charac erizing the high-pressure phases and the yielding process under shock loading.

INTRODUCTION

Understanding how nonhydrostatic stresses and high strain-rates can produce observed increases in structural phase transformation kinetics under shock heading is of fundamental, as well as provided, interest. Continuum level investigations of anock-induced transformations suffer from the inability to identify microscopic processes directly. Nonetheless, the full potential of continuum measurements has not been achieved because only longitudinal stress and particle velocity can be measured reliably, which leaves the deviatoric and mean stresses undetermined.

Measuring both shear and compression waves using embedded particle velocity ganges in comlined, compression and shear hading plate impact experiments augments continuum measurements with new information that is directly related to both the shear properties and mean stress of the shucked material.¹,³ The shear wave is a useful probe of the shear properties of the compressed state produced by the faster traveling compression wave. It is not of practical use for mapping the yield surface by varying the shear and compression wave amplitudes in successive experiments: The shear stress farmed to the compression wave is \sim 50% greater than that in the shear wave tassuming

a manageably large inclination of $0 \approx 20^{\circ}$ and Puisson ratio of 0.25). For a von Mises yield, this corresponds to the longitudinal stress at yield being

only 12% less than for uniaxial strain alone ($\theta=0^{\circ}$).

The efficacy of measuring both shear and compression waves for studying phase-transforming materials has been demonstrated recently in an investigation of polycrystalline CaCO₄, where it has been valuable for distinguishing effects of phase transformations from yielding and for characterizing both the high-pressure phases and the yielding process.^{3,4} Analogous results have been obtained or should be obtainable in non-phasetransforming materials. In contrast to previous results,^{4,2} in the phase-transforming material mean stress could not be directly computed from the experimentally-determined limb modulus values.⁴

EXPERIMENTAL METHOD

In initerial, electromagnetic gauges at selected depths mease, e transverse and longitudinal particle velocity histories associated with the propagation of large amplitude, one dimensional compression and shear waves. These waves are generated by parallel impact of two plates that are equally inclined to their direction of approach. The magnetic field and gauge orientations are chosen so that, under ideal conditions, the gauges respind to either longitudinal or transverse motion.⁵

Two experiments are needed at each impact velocity due narinly to the impracticality of placing more than four gauges in a sample.⁶ In the Inst, we simultaneously measure the tongitudinal and transverse motions with two gauges set at the same depth from the impact face. This measurement is important for correlating the shear and compression wave arrivals precisely when the latter has a complex structure due to yielding or structural phase transformations.⁵ In the duplicate experiment, we measure the longitudinal histories, only, at three depths to permit calculation of longitudinal stress and strain fields.

RESULTS

These measurements provide the following information at each impact velocity: Shock wave speed; sheac and longitudinal wave speeds at the perk compression; longitudinal stress longitudinal strain paths; shear stress shear strain paths. The shear and bulk moduli are obtained from the wave speeds and density. Measurements at a series of impact velocities provide the density variation of the moduli in the shocked state.

A series of compression shear experiments on Carrara marble samples at peak levels from 0.5 GPa to 6.3 GPa longitudinal stress demonstrate the possibility and value of measuring both alicar and longitudical waves in phase transforming materials.^{1,4} Mechanical yielding and two phase transformations occurred within this range. These processes prevented neither shear wave propagation nor the particle velocity measurements. Details of the response of CaCO3 are reported in [3] and [4]. Here we focus on two comparisons afforded by the shear wave measurements that help greatly in distinguishing phase transformations from yielding and in characterizing the high pressure phases and yielding process.

In Figure 1 the bulk modulus values in the shocked marble are plotted versus peak density compression. The open squares are estimates hased on the measured release wave speed.⁴ Because the elastic moduli as functions of density thor stress) represent the same physical quantities inider shock and hydrostatic loading the hilk modulus values perind a more discriminating conparison of material response under these loadure conditions than do wave speeds. The solid line represents the isentropic tork postably values de termined from affrasome wave speed measure ments of Oak Hall Intrestone 7. The two sets of broken lides give the isothering bulk modulis valacs det turned from hydrostatic compression curves to escore ervicil cale to 01/81/deshed) and

[9] (short dash). Values from a finite strain fit to the latter hydrostat are shown by the bold dashed line.⁴ The stability fields of the three phases of served under room temperature hydrostatic compression are indicated by the horizontal arrows.

The hulk modulus in the shocked marble shows the expected decrease due to changes in elastic properties neross the transformation to CaCO₃(11) and a subsequent increase that is comparable to that observed under hydrostatic loading up to the start of the CaCO₃(11) stability field. At 11% compression, the shocked marble bulk modulus lies for below the value determined from the hydrostat.⁴

Because the bulk modulus is rather insensitive to deformation or yielding, this discrepancy suggests the onset of a transformation to a phase other than CaCO₄(11) that has a substantially lower bulk modulus. The bulk modulus of compressed aragmite is not known, but it is the likely candidate for this new phase. This is supported by the calcite Hugomot crossing the CaCO₄(11) segment of the hydrostat by 14% compression and merging with the aragonite Hugomiot.⁴

In previous compression shear studies^{1,2} the mean stress in the shocked state was calculated from the density integral of the hulk modulus.



Engine 1. Bulk modulus to shocked Carrata nuclific (squares) versus tormalized density compared with hydrostatic data (see text)

Mean stress values are needed for separating deviatoric stress and strain-rate effects on materials response. Without them the usefulness of continuum shock wave measurements is curtailed substantially. Integrating the hulk modulus values shown in Figure 1 leads to unphysical negative stress deviators in the mixed-phase region. In phasetransforming materials the mean stress cannot be directly computed in this manner apparently because the experimentally-determined bulk modulus values in the mixed-phase region are frozen phase-composition values and so are not equal to the logarithmic density derivative of the Hugoniot mean stress.

The difficulty in calculating the Hugoniat mean stress can be understood by considering the schematics shown in Figure 2. In Figure 2a, the dushed curve is an idealized compression curve for a single crystal that undergoes a first-order phase transformation. The Hugoniot mean stress for a polycrystal of this same material is represented by the solid enrye. Figure 2b shows the corresponding bulk modulus density relations. The single crystal modulus is taken to decrease across the transformation, as is observed for calcite. The solid curve gives the relaxed polycrystal lalk modulus, K_{11x}. The relaxed and frozen modulus, K_{frz}, values at 3% compression are indicated by the dot and open circle, respectively; frozen modulus values are always intermediate between the pure component values.

The Hugoniot mean stress can be obtained by integrating the bulk modulus provided the latter corresponds to the slope of the mean stress-density curve. In the mixed phase region Kilx is systematically lower than K₁₀ because it includes the contribution from the volume change of the transtormed fraction of the sample. The speed of the, ideally, infinitesimal amplitude leading edge of the release wave is used to determine the hulk moduhis. As no change in phase composition occurs in the leading edge, the incastified wave speeds are trozen phase composition values. Consequently, integrating the experimentally determined hulk modulus density relation overestimates the mean stress and capilead to apparent strain softening or even negative stress deviators.

The shear wave amplitudes can provide information about nonterial strength that cannot be obtained from the wave speeds or the compression wave amplitudes. In phase transforming materials



Figure 2. Schematic of responses governed by frozen and relaxed bolk moduli.

this additional information can be important for distinguishing the effects of yielding from phase transformations. However, numerical simulations of the wave propagation using assumed material models are needed to analyze shear wave amplitudes.¹⁰ Nonetheless, it is readily apparent that, because shear wave amplitudes are related to material strength, the ability of the marble to support limite amplitude shear waves indicates that it retains significant strength ap to 6.3 GPa.⁴

Gupta's simulations illustrate the approach for analyzing shear wave amplitudes. The demonstrates that the shear wave is sensitive to the yield process during longitudinal unloading.^{10,11} Hence, combined with numerical simulations, the shear wave history during longitudinal unloading can be used to detect yielding and characterize portions of the yield surface.

Qualitative aspects of the marble strength response were inferred by comparing the results of Gupta's simulations for granite with the shear wave particle velocity histories in Carrara marble:¹¹ The marble yields due to passage of the compression wave, alone, at a longitudinal stress of 1.55 GPa or less: the marble exhibits yield tehavior that resembles a pressure-dependent response.

CONCLUSIONS

The investigation of Carrara marble demonstrates that shear waves can be measured in a material undergoing shock-induced phase transformations and that these measurements provide additional information not obtainable from aniaxial strain, plate impact experiments. The shear and longitudinal wave speeds, together with calculated peak densities,¹² provide shear and bulk moduli values of the high-pressure phases. In mixedphase regions the moduli so determined are frozen phase-composition quantities.

The Hugoniot mean stress cannot be directly computed from these link modulus values. There is currently no proven method for determining the Hugoniot mean stress in an arbitrary solid. Lateral stress gauge measurements might eventually provide this crucial capability (in this regard see M. K. W. Wong in these Proceedings). Nonetheless, the bulk modulus data provide strong additional constraints on material models.

Shear wave amplitudes can be useful for characterizing inelastic response, hut they do not permit simple determination of the yield surface.

Together the shear and compression wave measurements more fully characterized the shocked state, but they were insufficient for answering many specific questions about the dynamic response of calcite rocks, even on the continuum level.⁴ It was heneficial to the investigation that the first phase transformation was rapid, easily identifiable, and occurred wittin a stress range that was easy to study.

*This work was performed at Washington State University accollaboration with Y. M. Gupta.

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